

Vertical motion observed with GPS:

What can we learn about regional geophysical signals, Earth structure, and rheology?

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Introduction

Tectonic studies utilizing observations of surface displacements from GPS are based predominantly on observations of horizontal displacements, while vertical displacements mostly provide constraints for loading-related geophysical processes. A main reason for this disjoint use of GPS observations is that tectonic signals often are larger in the horizontal component while loading signals dominantly are contained in the vertical component. Moreover, the horizontal components of GPS-determined displacements have a precision about three times that of the vertical component.

Considering the growing number of PBO stations with records potentially long enough to determine reliable secular trends, it is worthwhile to ask the question to what extent vertical displacement time series can be used to constrain tectonic processes, Earth structure, and rheology. In order to answer this question, we study the anatomy of time series of vertical displacements determined from regional and local GPS networks in North America and derive the spatio-temporal pattern of vertical crustal motion as seen by these networks.

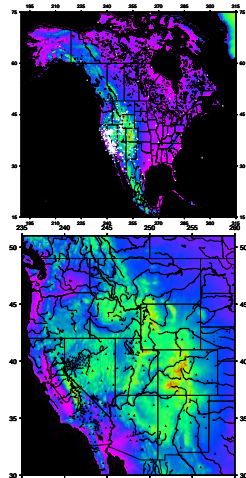


Figure 1: GPS stations used in the present analysis. Upper diagram: North America. Lower diagram: GPS station distribution in the Western U.S.A.

The Data

We have collected GPS observations for North American stations from the publicly available archives at UNAVCO, SOPAC, and CORS as well as our own MAGNET archive (Figure 1). The focus of our studies is the South-Western part of the U.S., and there station distribution is densest (Figure 2).

All the GPS data were homogeneously processed with the GIPSY-OASIS II (GIPSY) software package of Jet Propulsion Laboratory (JPL) using the *Precise Point Positioning* (PPP) method (Zumberge et al., 1997) to determine daily coordinates, with ambiguity resolution applied across the entire network by automatic selection of the ionospheric- or pseudorange-widelane method (Blewitt, 1989). Satellite orbit and clock parameters were provided by JPL. Ionosphere-free combinations of carrier phase and pseudorange were processed every 5 minutes. To account for atmosphere effects that can bias the inference of the antenna phase center height we solve simultaneously for station height, zenith path delay and tropospheric gradient parameters. We processed the data with a satellite elevation cut-off angle of 15° in order to minimize the effects of atmospheric refraction and multipath. Compared to lower cutoff elevations, this results, on the one hand, in less data used and hence more uncertainty in individual daily positions, and, on the other hand, in lower bias of station coordinates associated with atmospheric refraction of the GPS signal, which is greatest for satellites at low elevation. Ambiguity resolution was carried out on the complete North American network.

All time series were determined in ITRF2000. Time series of daily station displacements are computed with respect to the coordinates at epoch 2000/02/01. No spatial filtering has been applied to the time series. These time series were then used to determine vertical rates taking into account semi-annual and annual harmonic constituents. For time series analysis, the time window from 2000/02/01 to 2006/10/31 was used in order to avoid offsets due to the transition of JPL from ITRF2000 to ITRF2005.

Time series

Figure 2 shows time series of vertical displacements for selected North American stations. Day-to-day variation and intraseasonal to interannual variations are of the order of ±20 mm and ±10 mm, respectively. Most station exhibit a linear trend, and a seasonal cycle is discernible. For stations separated on the order of 10³ km, considerable differences in the intraseasonal and interannual variations are visible (see Figure 2, left diagrams). On local scale (up to 500 km, see the example of the Basin and Range Province, Figure 2), the differences on intra- to interannual time scales are very small, while the main difference are in the day-to-day variations (Figure 2, right diagrams).

The displacement time series are a combination of geophysical signals and apparent displacements induced by the GPS system and data analysis. Concerning the observed interannual variations, a key problem is the separation of instrumental and analysis effects from the geophysical signals. The interannual variations have surprisingly small spatial scales, of the order of 10³ km (see below), with significant spatial variations over North America. If these variations were solely due to GPS-system and analysis effects, we would expect larger spatial scales in these variations.

Spectrum

The discrete part of the spectrum of the daily displacement time series is generally dominated by a semi-annual and annual harmonic constituent, denoted here as Sa and Ssa, respectively (Figure 3). The ratio of the amplitudes of Sa and Ssa changes from station to station.

Cross-correlation

The cross-correlation coefficient for the vertical displacement time series (station pairs) varies considerably between ±0.8 even for nearby stations. (Figure 4, left diagram). No clear spatial pattern is discernible. However, if linear trend and annual and semi-annual constituents are removed, cross-correlation for the residual time series increases considerably (Figure 4, right diagram). Taking individual stations as base, even for distances of order 10³ km, correlation coefficients are generally between 0.6 and 0.8 (Figure 5). For distances larger than order 10³ km, the cross-correlation decreases and breaks down for distances larger than 2 · 10³ km. This latter feature is somewhat uncertain due to the sparse number of station pairs with great distances.

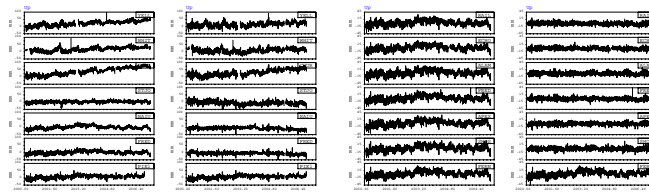


Figure 2: Time series of vertical displacements. Left two diagrams: Stations distributed from 31° to 65°N. Right two diagrams: Stations in the Basin and Range Province. In each pair, the time series of vertical displacements are shown in the left diagram, while the right diagram shows for each station the difference to the time series at the bottom. The offset around Nov. 5, 2006 is due to the transition of JPL to ITRF2005.

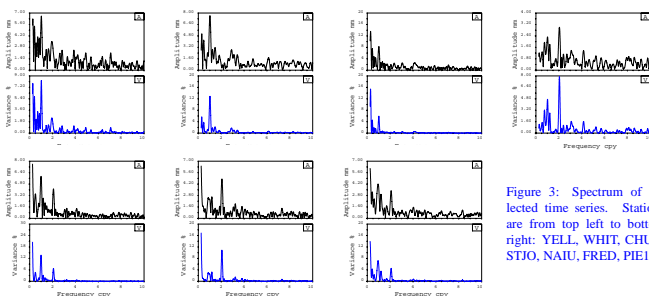


Figure 3: Spectrum of selected time series. Stations are from top left to bottom right: YELL, WHIT, CHUR, STJO, NAIU, FRED, PIE1.

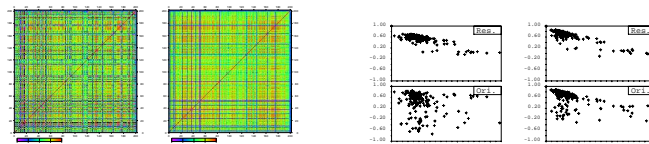


Figure 4: Cross-correlation of time series of vertical displacements. Stations are approximately order from northwest to southeast. Left: Correlation of original time series. Right: Correlation of residual time series after trend and seasonal cycle have been removed.

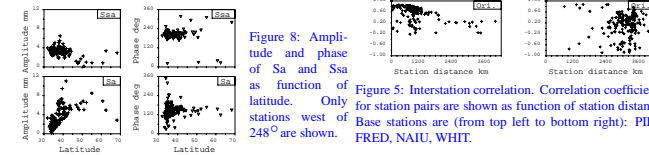


Figure 5: Amplitude and phase of Sa and Ssa as function of latitude. Only stations west of 248° are shown.

Secular Trends.

Generally, at interannual periods, the GPS time series of vertical motion show significant nonlinear components potentially contaminating the derived secular velocity field (see Figure 2). Determining reliable secular trends for time series with variations at interannual to decadal and longer time scales requires time series longer than these scale. This is well known from studies of climatological and oceanographic time series, where often 20 to 50 years are considered the minimum record length for trend estimation. However, in the case of climate and oceanography, the signals are mainly due to the Earth system processes, which can be modeled to a large extent. For vertical GPS time series, a significant portion may come from the observing system and data analysis, thus complicating the situation considerably. Nevertheless, the spatial pattern of the secular vertical velocity field displays features potentially related to geophysical processes, including tectonics (Figure 6). In particular, the Basin and Range Province shows mainly significant uplift, which is contrary to the expectation of an extensional basin driven by gravitational potential energy.

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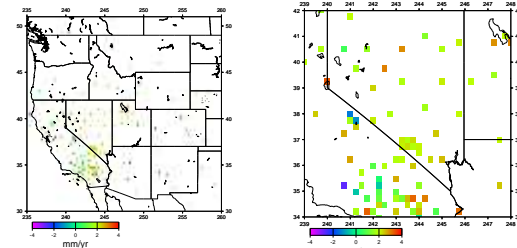


Figure 6: Vertical rates in the Western U.S. Left: Western U.S., right: Basin and Range Province.

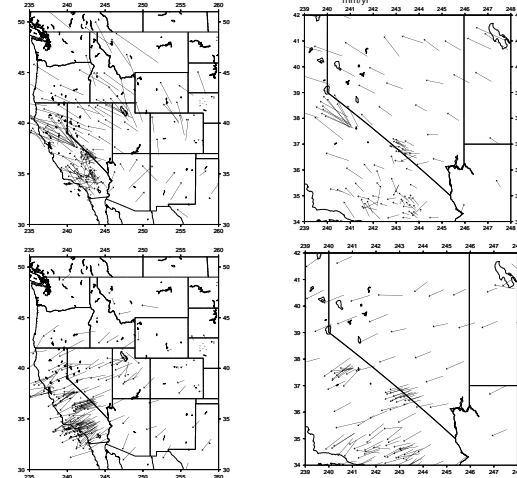


Figure 7: Seasonal cycle in vertical displacements. Left: Western U.S., right: Basin and Range Province. Top: Annual harmonic constituent. Bottom: Semi-annual constituent.

Seasonal Cycle

Sa and Ssa show high spatial coherency in amplitude and phase over large geographical areas except for few deviations which most likely are due to local influences (Figure 7). The spatial coherency underlines that GPS picks up geophysical signals at the seasonal time scale. The amplitudes of Sa increases from South to North up to latitude 50°N and is smaller again at high latitudes (Figure 8). For Ssa, a general decrease of the amplitude from South to North is found. The spatial pattern of Sa and Ssa as seen by GPS can be used to identify the forcing geophysical signals (loading) and separate these from potential system-induced contribution.

Conclusions

Beside spatially homogeneous annual and semi-annual harmonic constituents the GPS vertical displacement time series have significant interannual variations with spatial scales of a few hundred to a thousand kilometer, potentially biasing the secular trends determined from time series of a few years to decades. Strictly speaking, if we do not understand the interannual variations of GPS vertical time series, we cannot safely interpret the secular signal in terms of global and regional processes. Spatial filtering, which has been successfully applied to extract the secular signal from horizontal components (Wdowinski et al., 2004, e.g.) does not help to eliminate the interannual signals in the vertical. Part of these variations are due to reference frame instability. Therefore, improving the reference frame seems to be a prerequisite for the interpretation of vertical motion in terms of large-scale tectonic process and must have highest priority. A significant improvement is expected by modeling load-induced deformations in the reference frame determination. The geophysical loading models can be validated against the spatial patterns of Sa and Ssa as seen by GPS.

References

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